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14. ABSTRACT The HYbrid Coordinate Ocean Model (HYCOM) is isopycnal in the open, stratified ocean, but uses the layered continuity equation to make a dynamically smooth transition to a terrain-following coordinate in shallow coastal regions, and to z-level coordinates in the mixed layer and/or unstratified seas. The hybrid coordinate extends the geographic range of applicability of traditional isopycnic coordinate circulation models toward shallow coastal seas and unstratified parts of the world ocean. The principal goal of this DoD Challenge project is to provide a near real time depiction of the three-dimensional global ocean state at fine resolution. On a 7 km (mid-latitude) grid, HYCOM has sufficient resolution to provide a baseline depiction of the most of the world's coastal areas, and will be the highest horizontal resolution world's wide source for realistic offshore boundary conditions for realistic offshore boundary conditons for nested coastal models.					
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# Global Ocean Prediction Using HYCOM

Alan J. Wallcraft, E. Joseph Metzger, and Harley E. Hurlburt  
*US Naval Research Laboratory (NRL-SSC), Stennis Space Center, MS*  
{wallcraft, metzger, hurlburt}@nrlssc.navy.mil

Eric P. Chassignet and Zulema D. Garraffo  
*Rosenstiel School of Marine and Atmospheric Science,  
University of Miami, Miami, FL*  
{echassignet, zgarraffo}@rsmas.miami.edu

Ole Martin Smedstad  
*Planning Systems, Inc., Slidell, LA*  
smedstad@nrlssc.navy.mil

## 1. Introduction

The HYbrid Coordinate Ocean Model (HYCOM) (Bleck, 2002) is isopycnal in the open, stratified ocean, but uses the layered continuity equation to make a dynamically smooth transition to a terrain-following coordinate in shallow coastal regions, and to z-level coordinates in the mixed layer and/or unstratified seas. The hybrid coordinate extends the geographic range of applicability of traditional isopycnic coordinate circulation models, such as NLOM and MICOM, toward shallow coastal seas and unstratified parts of the world ocean. It maintains the significant advantages of an isopycnal model in stratified regions while allowing more vertical resolution near the surface and in shallow coastal areas, hence providing a better representation of the upper ocean physics. HYCOM is designed to provide a major advance over the existing operational and preoperational global ocean prediction systems, since it overcomes design limitations of the present systems as well as limitations in vertical and horizontal resolution. The result should be a more streamlined system with improved performance and an extended range of applicability (e.g., the present systems are seriously limited in shallow water and in handling the transition from deep to shallow water).

The principal goal of this DoD Challenge project is to provide a near real time depiction of the three-dimensional global ocean state at fine resolution ( $1/12^\circ$  on the equator,  $\sim 7$  km at mid-latitudes, and  $\sim 4$  km in the Arctic). This HYCOM-based system will include an embedded ice model and the capability to host nested littoral models with even higher resolution. It will be the next generation eddy-resolving operational global ocean nowcast/forecast system at NAVOCEANO, with transition from Research and Development to

NAVOCEANO planned for 2007. The resolution should increase to  $1/25^\circ$  ( $\sim 3-4$  km at mid-latitudes) by the end of the decade. Starting in mid-2006, a major sub-goal of this effort is participation, using  $1/12^\circ$  global HYCOM, in the multinational Global Ocean Data Assimilation Experiment (GODAE). It is designed to help justify a permanent operational global ocean observing system by demonstrating useful real-time global ocean products with a customer base.

On a 7 km (midlatitude) grid, HYCOM has sufficient resolution to provide a baseline depiction of the most of the world's coastal areas, and will be the highest horizontal resolution world wide source for realistic offshore boundary conditions for nested coastal models. Other applications for the models and the nowcast/forecast systems include assimilation and synthesis of global satellite surface data; ocean prediction; optimum track ship routing; search and rescue; antisubmarine warfare and surveillance; tactical planning; sea surface temperature for long range weather prediction; inputs to shipboard environmental products; environmental simulation and synthetic environments; observing system simulations; ocean research; inputs to biogeochemical and optical models; pollution and tracer tracking and inputs to water quality assessment.

## 2. Model Configuration

There are several projections that allow the Arctic to be included in a global ocean model by moving the singularity at the pole over land. For the HYCOM global configuration, we use an Arctic dipole patch matched to a standard Mercator grid at  $47^\circ\text{N}$ . Unlike most other poleshifting projections, this has the advantage that all grid points below  $47^\circ\text{N}$  are unchanged. Since HYCOM supports general orthogonal curvilinear grids, this

requires no changes to the standard model code and array structure except a special halo exchange at the northern edge of the logically rectangular domain (Figure 1). Locating the dipoles at 47°N gives good resolution in the Arctic Ocean (7 km at midlatitude vs. 3.5 km at the North Pole), where the radius of deformation is small. For our target resolution (1/12° at the equator), the array size is 4500 by 3298 with 26–32 hybrid layers in the vertical. The complete system will include the Los Alamos CICE seaice model (Hunke and Lipscomb, 2004) on the same grid. The ocean and ice models will run simultaneously, but on separate sets of processors, communicating via an Earth System Modeling Framework (ESMF) based coupler (Hill et al., 2004). A typical configuration would use ~750 processors for the HYCOM component plus a much smaller number for CICE (since it does not need to be run on the ice free ocean), chosen to make the two run in the same amount of wall time.

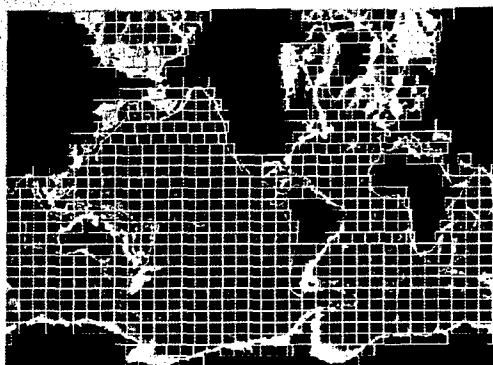


Figure 1. 1/12° global domain with 36×32 tiling, and 781 MPI tasks. The Arctic “wraps” between the left and right halves of the top edge.

The basic parallelization strategy is domain decomposition, i.e., the region is divided up into smaller subdomains, or tiles, and each processor “owns” one tile. Figure 1 shows the tiling we are currently using for the 1/12° global domain, consisting of 1152 (36 by 32) approximately equalsized tiles, but 371 “all land” tiles are discarded leaving 781 MPI tasks. For more details on scalability see the companion paper on our Capability Applications Project (Wallcraft et al., 2005).

### 3. Progress

In the first year of this Challenge project, we will primarily be running nonassimilative simulations. These are initialized from an oceanic climatology of temperature and salinity and forced by climatological atmospheric wind and thermal fields. After initialization, the only oceanic data source is relaxation to climatological surface salinity. Figure 2 shows sea surface height and temperature for January 1<sup>st</sup>; ice coverage is in gray.

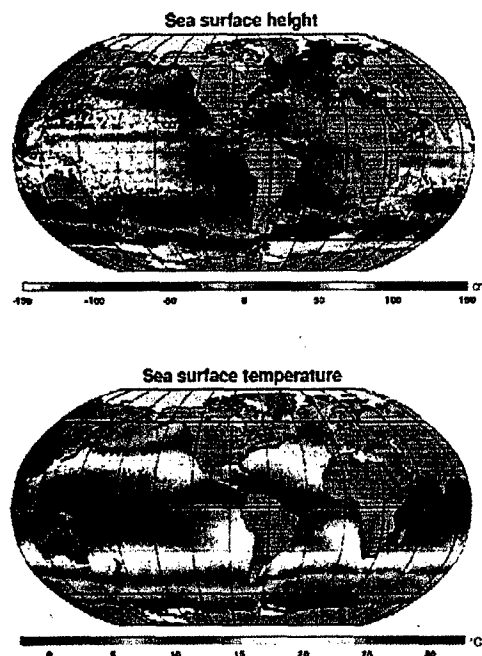


Figure 2. Sea surface height and temperature, and ice coverage (gray), from 1/12° global HYCOM on January 1<sup>st</sup>

Our goal is to couple the sea ice and ocean models (CICE and HYCOM), but our simulations to date are using the simple thermodynamic sea ice model that is included with HYCOM. We have CICE running standalone on our global arctic patch grid (for the Arctic only), but are waiting for ESMF-based coupling between HYCOM and CICE to be available before running a coupled case. Even without ice dynamics, the seasonal cycle of ice coverage is good overall. Figure 3 compares the total northern and southern hemispheric ice area to the NOAA OISST climatology (Reynolds et al., 2002). One reason for the good agreement is that the atmospheric forcing is based on an accurate ice extent and this provides a strong tendency for the ocean/seaice system to form ice appropriately.

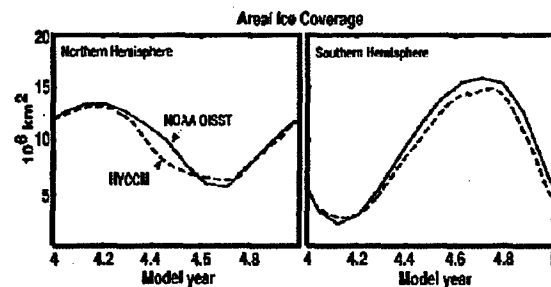


Figure 3. Ice coverage vs. time from 1/12° global HYCOM (dashed) and NOAA OISST climatology (solid) for the northern (left) and southern (right) hemispheres

Figure 4 compares HYCOM's annual mean zonal velocity along the equator in the Pacific with buoybased observations from the Tropical Atmosphere Ocean (TAO) project (Johnson et al., 2002). The Equatorial Undercurrent (EUC) is well developed in HYCOM, with a maximum of 90 cm/s vs. 100 cm/s for the observations. The model also reproduces the EUC's west to east shoaling and strengthening.

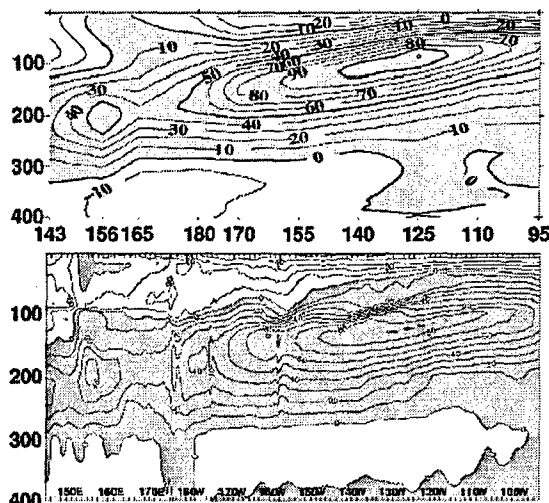


Figure 4. Annual mean zonal velocity (positive eastward, in gray) vs. depth along the equator in the Pacific, from (top) 10 Tropical Atmosphere Ocean (TAO) moorings at the labeled latitudes, and (bottom) 1/12° global HYCOM. Contour interval is 10 cm/s.

#### 4. Plan

The majority of the first year simulations will be freerunning, with atmospheric forcing only. In the second year we will concentrate on adding ocean data assimilation and, starting in mid 2006, we will perform a nowcast every day and a 30-day forecast every week in near real time.

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